



Understanding the Science of Climate Change

Talking Points: Impacts to Western Mountains and Forests

Natural Resource Report NPS/NRPC/NRR—2009/090



ON THE COVER

Heaven's Peak, Glacier National Park; National Park Service photo.

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I. Introduction

Purpose

Climate change presents significant risks to our nation's natural and cultural resources. Although climate change was once believed to be a future problem, there is now unequivocal scientific evidence that our planet's climate system is warming (IPCC 2007a). While many people understand that human emissions of greenhouse gases have caused recent observed climate changes, fewer are aware of the specific impacts these changes will bring. This document is part of a series of bio-regional summaries that provide key scientific findings about climate changes in and impacts to protected areas. The information is intended to provide a basic understanding of the science of climate change, known and expected impacts to resources and visitor experience, and actions that can be taken to mitigate and adapt to change. The statements may be used to communicate with managers, frame interpretive programs, and answer general questions to the public and the media. They also provide helpful information to consider in the developing sustainability strategies and long-term management plans.

Audience

The Talking Points are primarily intended to provide park and refuge area managers and staff with accessible, up-to-date information about climate change and impacts to the resources they protect.

Organizational Structure

Following the Introduction are three major sections of the document: a general section on Global Climate Change, a Regional section that provides information on change to western mountains and forests, and a section outlining No Regrets Actions that can be taken now to mitigate and adapt to change. Within the global and regional sections, scientific information is organized in two basic ways: 1) the type of change or impact and 2) the general level of uncertainty in the statement.

1) Type of Change or Impact

Global Section

- A. Temperature and Greenhouse Gases
- B. Water, Snow, and Ice
- C. Vegetation and Wildlife
- D. Hazards

Regional Section

- A. Temperature
- B. The Water Cycle (including snow, ice, lake levels, and sea level)
- C. Vegetation (plant cover, species range shifts, and phenology)
- D. Wildlife (aquatic, marine, and terrestrial animals,

range shifts, invasive species, migration, phenology)

E. Disturbance (including invasive species, plant pests and pathogens, fire, flooding, and erosion)

F. Visitor Experience

2) Level of uncertainty

Information contained in this document is derived from the published results of a range of scientific research including historical data, empirical (observed) evidence, and model projections (which may use observed or theoretical relationships). While all of the statements are informed by science, not all statements carry the same level of confidence or scientific certainty. Identifying uncertainty is an important part of science but can be a major source of confusion for decision makers and the public. In the strictest sense, all scientific results carry some level of uncertainty because the scientific method can only "prove" a hypothesis to be false. However, in a practical world, society routinely elects to make choices and select options for actions that carry an array of uncertain outcomes.

The statements in this document have been organized to help managers and their staffs differentiate among current levels of uncertainty in climate change science. In doing so, the document aims to be consistent with the language and approach taken in the Fourth Assessment on Climate Change reports by the Intergovernmental Panel on Climate Change. However, this document discriminates among only three different levels of uncertainty and does not attempt to ascribe a specific probability to any particular level. These are qualitative rather than quantitative categories and are based on the following:

- "What scientists know" are statements based on measurable data and historical records. These are statements for which scientists generally have high confidence and agreement because they are based on actual measurements and observations. Events under this category have already happened or are very likely to happen in the future.
- "What scientists think is likely" represents statements beyond simple facts; these are derived from some level of reasoning or critical thinking. They result from projected trends, well tested climate or ecosystem models, or empirically observed relationships (statistical comparisons using existing data).
- "What scientists think is possible" are statements that use a higher degree of inference or deduction than the previous categories. These are based on research about processes that are less well understood, often involving dynamic interactions among climate and complex ecosystems. However, in some cases, these statements represent potential future conditions of greatest concern, because they may carry the greatest risk to protected area resources.

II. Global Climate Change

Unless otherwise noted, all statements and figures in this section were derived from the Intergovernmental Panel on Climate Change: Synthesis Report (IPCC 2007a).

The IPCC is a scientific intergovernmental, international body established by the World Meteorological Organization (WMO) and by the United Nations Environment Programme (UNEP). The information the IPCC provides in its reports is based on scientific evidence and reflects existing consensus viewpoints within the scientific community. The comprehensiveness of the scientific content is achieved through contributions from experts in all regions of the world and all relevant disciplines including, where appropriately documented, industry literature and traditional practices, and a two stage review process by experts and governments.

Definition of climate change: The IPCC defines climate change as a change in the state of the climate that can be identified (e.g. using statistical tests) by changes in the mean and/or the variability of its properties, and that persists for an extended period, typically decades or longer.

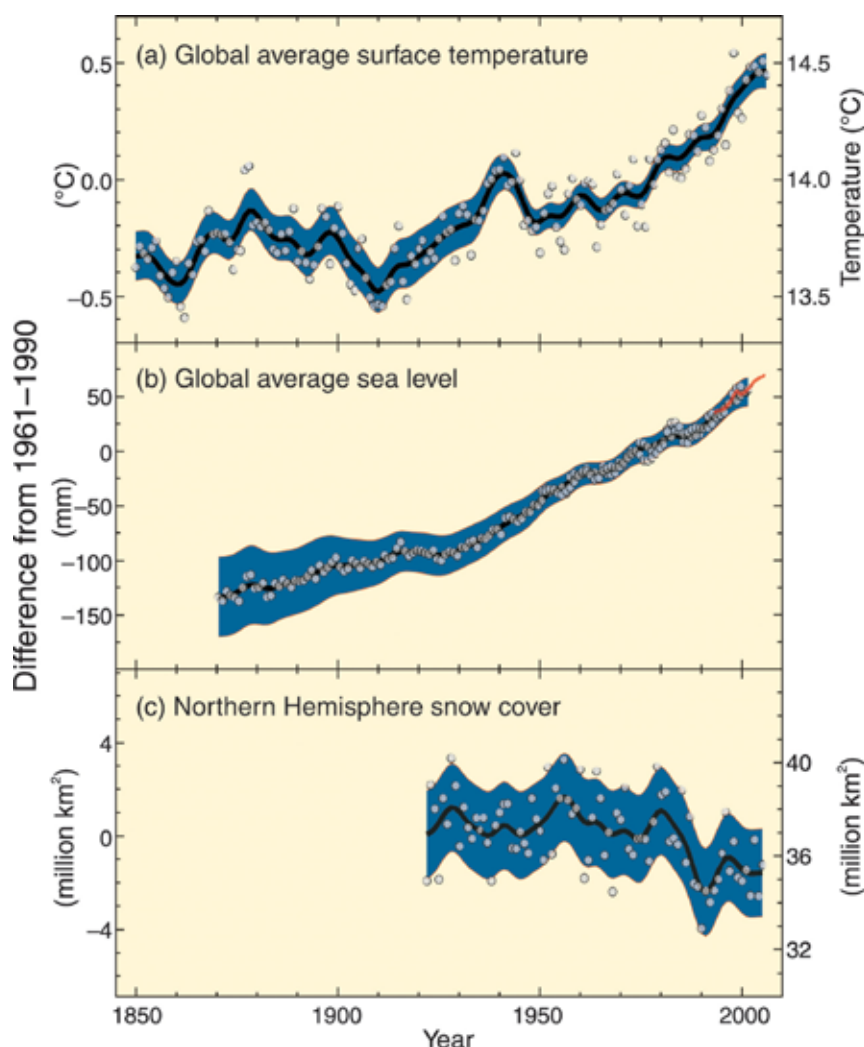


Figure 1. Observed changes in (a) global average surface temperature; (b) global average sea level from tide gauge (blue) and satellite (red) data and (c) Northern Hemisphere snow cover for March-April. All differences are relative to corresponding averages for the period 1961-1990. Smoothed curves represent decadal averaged values while circles show yearly values. The shaded areas are the uncertainty intervals estimated from a comprehensive analysis of known uncertainties (a and b) and from the time series (c) (IPCC 2007).

A. Temperature and Greenhouse Gases

What scientists know...

- Warming of the Earth's climate system is unequivocal, as evidenced from increased air and ocean temperatures, widespread melting of snow and ice, and rising global average sea level (Figure 1).
- In the last 100 years, global average surface temperature has risen about 0.74°C over the previous 100-year period, and the rate of warming has doubled from the previous century. Eleven of the 12 warmest years in the instrumental record of global surface temperature since 1850 have occurred since 1995 (Figure 1).
- Although most regions over the globe have experienced warming, there are regional variations: land regions have warmed faster than oceans and high northern latitudes have warmed faster than the tropics. Average Arctic temperatures have increased at almost twice the global rate in the past 100 years, primarily because loss of snow and ice results in a positive feedback via increased absorption of sunlight by ocean waters (Figure 2).
- Over the past 50 years widespread changes in extreme temperatures have been observed, including a decrease in cold

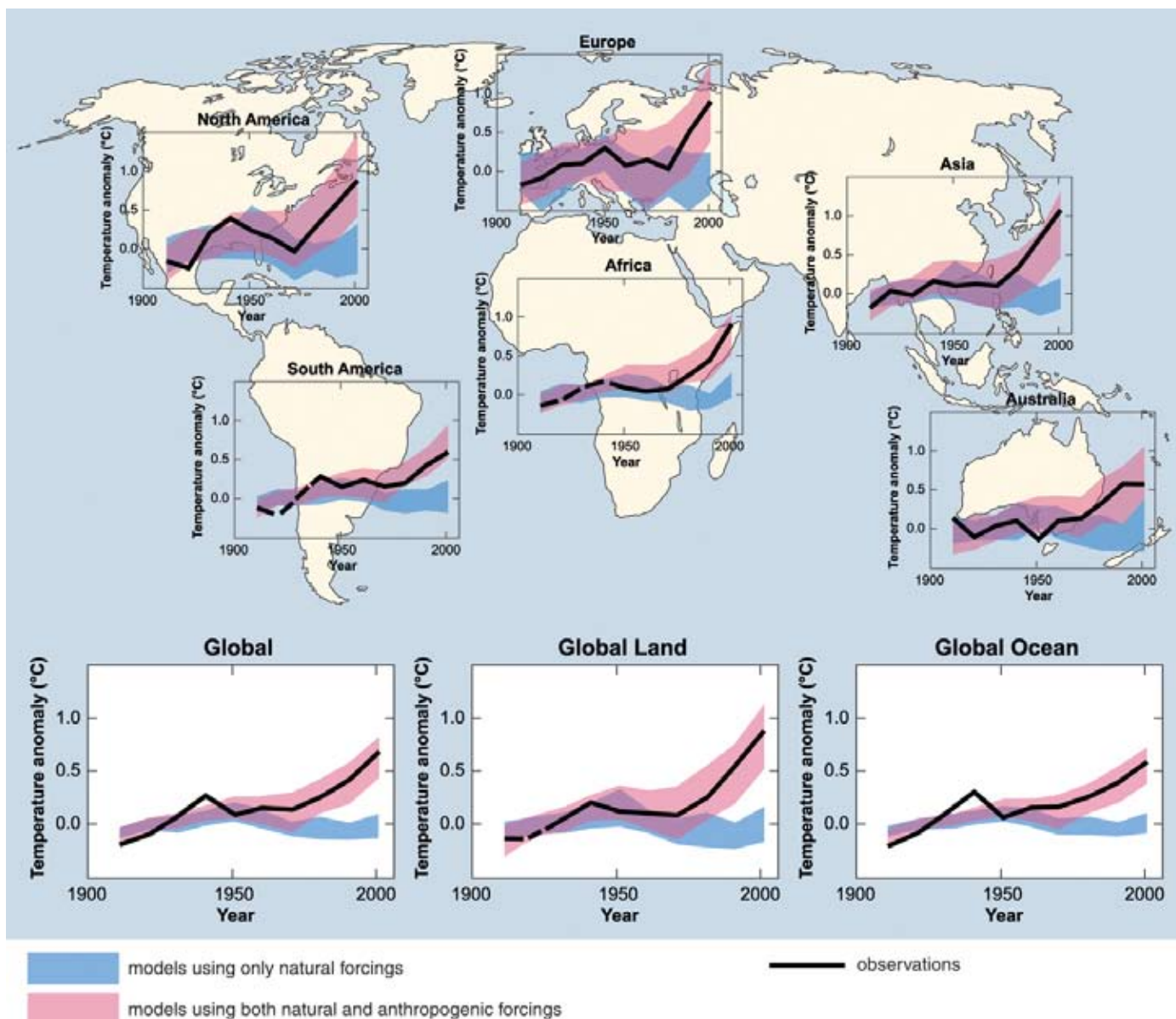


Figure 2. Comparison of observed continental- and global-scale changes in surface temperature with results simulated by climate models using either natural or both natural and anthropogenic forcings. Decadal averages of observations are shown for the period 1906-2005 (black line) plotted against the centre of the decade and relative to the corresponding average for the period 1901-1950. Lines are dashed where spatial coverage is less than 50%. Blue shaded bands show the 5 to 95% range for 19 simulations from five climate models using only the natural forcings due to solar activity and volcanoes. Red shaded bands show the 5 to 95% range for 58 simulations from 14 climate models using both natural and anthropogenic forcings (IPCC 2007).

days and nights and an increase in the frequency of hot days, hot nights, and heat waves.

- Winter temperatures are increasing more rapidly than summer temperatures, particularly in the northern hemisphere, and there has been an increase in the length of the frost-free period in mid- and high-latitude regions of both hemispheres.
- Climate change is caused by alterations in the energy balance within the atmosphere and at the Earth's surface. Factors that affect Earth's energy balance are the atmospheric concentrations of greenhouse gases and aerosols, land surface properties, and solar radiation.
- Global atmospheric concentrations of greenhouse gases have increased significantly since 1750 as the result of human activities. The principal greenhouse gases are carbon dioxide (CO₂), primarily from fossil fuel use and land-use change; methane (CH₄) and nitrous oxide (N₂O), primarily from agriculture; and halocarbons (a group of gases containing fluorine, chlorine or bromine), principally engineered chemicals that do not occur naturally.
- Direct measurements of gases trapped in ice cores demonstrate that current CO₂ and CH₄ concentrations far exceed the natural range over the last 650,000 years and have increased markedly (35% and 148% respectively), since the beginning of the industrial era in 1750.

- Both past and future anthropogenic CO₂ emissions will continue to contribute to warming and sea level rise for more than a millennium, due to the time scales required for the removal of the gas from the atmosphere.
- Warming temperatures reduce oceanic uptake of atmospheric CO₂, increasing the fraction of anthropogenic emissions remaining in the atmosphere. This positive carbon cycle feedback results in increasingly greater accumulation of atmospheric CO₂ and subsequently greater warming trends than would otherwise be present in the absence of a feedback relationship.
- There is very high confidence that the global average net effect of human activities since 1750 has been one of warming.
- Scientific evidence shows that major and widespread climate changes have occurred with startling speed. For example, roughly half the north Atlantic warming during the last 20,000 years was achieved in only a decade, and it was accompanied by significant climatic changes across most of the globe (NRC 2002).
- than during any other 50-year period in the last 500 years and likely the highest in at least the past 1300 years.
- Most of the warming that has occurred since the mid-20th century is very likely due to increases in anthropogenic greenhouse gas concentrations. Furthermore, it is extremely likely that global changes observed in the past 50 years can only be explained with external (anthropogenic) forcings (Figure 2).
- There is much evidence and scientific consensus that greenhouse gas emissions will continue to grow under current climate change mitigation policies and development practices. For the next two decades a warming of about 0.2°C per decade is projected for a range of emissions scenarios; afterwards, temperature projections increasingly depend on specific emissions scenarios (Table 1).
- It is very likely that continued greenhouse gas emissions at or above the current rate will cause further warming and result in changes in the global climate system that will be larger than those observed during the 20th century.

What scientists think is likely...

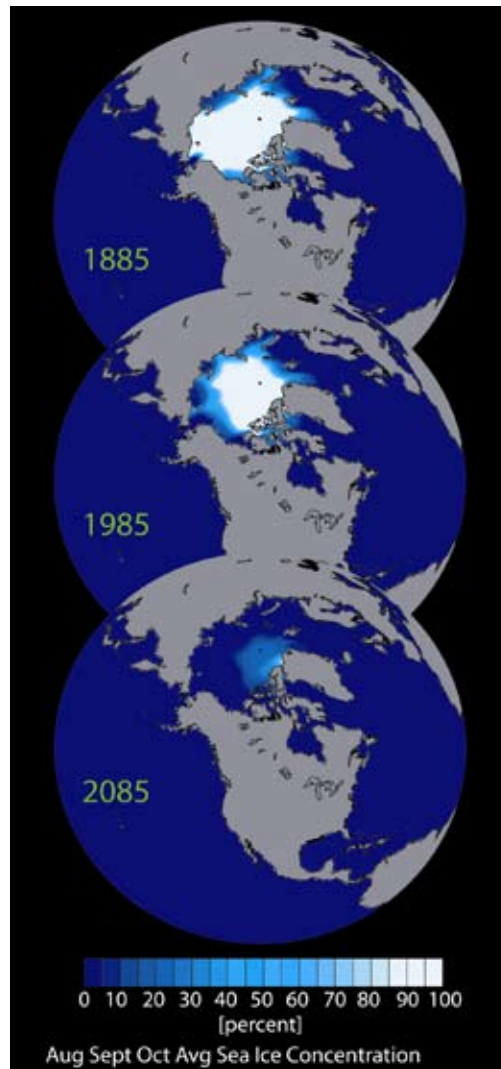
- Anthropogenic warming over the last three decades has likely had a discernible influence at the global scale on observed changes in many physical and biological systems.
- Average temperatures in the Northern Hemisphere during the second half of the 20th century were very likely higher
- It is very likely that hot extremes, heat waves and heavy precipitation events will become more frequent. As with current trends, warming is expected to be greatest over land and at most high northern latitudes, and least over the Southern Ocean (near Antarctica) and the northern North Atlantic Ocean.

Table 1. Projected global average surface warming at the end of the 21st century, adapted from (IPCC 2007b).

Notes: a) Temperatures are assessed best estimates and likely uncertainty ranges from a hierarchy of models of varying complexity as well as observational constraints. b) Temperature changes are expressed as the difference from the period 1980-1999. To express the change relative to the period 1850-1899 add 0.5°C. c) Year 2000 constant composition is derived from Atmosphere-Ocean General Circulation Models (AOGCMs) only.

Emissions Scenario	Temperature change(°C at 2090-2099 relative to 1980-1999) ^{a, b}	
	Best estimate	Likely range
Constant year 2000 concentrations ^a	0.6	0.3 - 0.9
B1 scenario	1.8	1.1 - 2.9
B2 scenario	2.4	1.4 - 3.8
A1B scenario	2.8	1.7 - 4.4
A2 scenario	3.4	2.0 - 5.4
A1F1 scenario	4.0	2.4 - 6.4

Figure 3. Sea ice concentrations (the amount of ice in a given area) simulated by the GFDL CM2.1 global coupled climate model averaged over August, September and October (the months when Arctic sea ice concentrations generally are at a minimum). Three years (1885, 1985 & 2085) are shown to illustrate the model-simulated trend. A dramatic reduction of summertime sea ice is projected, with the rate of decrease being greatest during the 21st century portion. The colors range from dark blue (ice free) to white (100% sea ice covered); Image courtesy of NOAA GFDL.



What scientists think is possible...

- Global temperatures are projected to increase in the future, and the magnitude of temperature change depends on specific emissions scenarios, and ranges from a 1.1°C to 6.4°C increase by 2100 (Table 1).
- Anthropogenic warming could lead to changes in the global system that are abrupt and irreversible, depending on the rate and magnitude of climate change.
- Roughly 20-30% of species around the globe could become extinct if global average temperatures increase by 2 to 3°C over pre-industrial levels.

B. Water, Snow, and Ice

What scientists know...

- Many natural systems are already being affected by increased temperatures, particularly those related to snow, ice, and frozen ground. Examples are decreases in snow and ice extent, especially of mountain glaciers; enlargement and increased numbers of glacial lakes; decreased permafrost extent; increasing ground instability in permafrost regions and rock avalanches in mountain regions; and thinner sea ice and shorter freezing seasons of lake and river ice (Figure 3).

- Annual average Arctic sea ice extent has shrunk by 2.7% per decade since 1978, and the summer ice extent has decreased by 7.4% per decade. Sea ice extent during the 2007 melt season plummeted to the lowest levels since satellite measurements began in 1979, and at the end of the melt season September 2007 sea ice was 39% below the long-term (1979-2000) average (NSIDC 2008) (Figure 4).

- Global average sea level rose at an average rate of 1.8 mm per year from 1961 to 2003 and at an average rate of 3.1 mm per year from 1993 to 2003. Increases in sea level since 1993 are the result of the following contributions: thermal expansion, 57%; melting glaciers and ice caps, 28%, melting polar ice sheets, 15%.

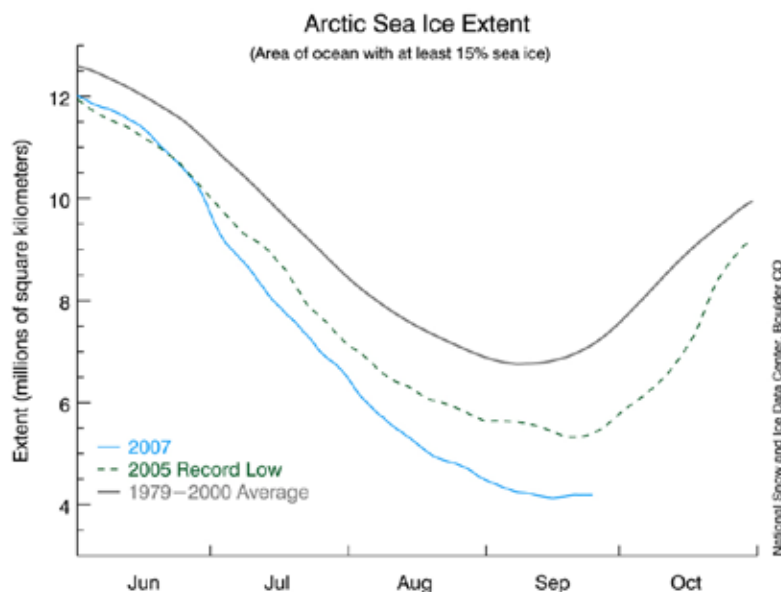


Figure 4. Arctic sea ice in September 2007 (blue line) is far below the previous low record year of 2005 (dashed line), and was 39% below where we would expect to be in an average year (solid gray line). Average September sea ice extent from 1979 to 2000 was 7.04 million square kilometers. The climatological minimum from 1979 to 2000 was 6.74 million square kilometers (NSIDC 2008).

- During the 20th century increasing atmospheric CO₂ has driven an increase in the average global ocean temperature of 0.74°C, and has reduced seawater carbonate concentrations to levels that may compromise carbonate accretion in marine organisms and limit formation of coral reefs (Hoegh-Guldberg et al. 2007).
- The oceanic uptake of anthropogenic carbon since 1750 has resulted in increased ocean acidification (average decrease in pH of 0.1 units), with observed and potential negative consequences for marine ecosystems.
- Even if greenhouse gas concentrations are stabilized at current levels thermal expansion of ocean waters (and resulting sea level rise) will continue for many centuries, due to the time required to transport heat into the deep ocean.
- Observations since 1961 show that the average global ocean temperature has increased to depths of at least 3000 meters, and that the ocean has been taking up over 80% of the heat added to the climate system.
- Hydrologic effects of climate change include increased runoff and earlier spring peak discharge in many glacier- and snow-fed rivers, and warming of lakes and rivers.
- Runoff is projected to increase by 10 to 40% by mid-century at higher latitudes and in some wet tropical areas, and to decrease by 10 to 30% over some dry regions at mid-latitudes and dry tropics. Areas in which runoff is projected to decline face a reduction in the value of the services provided by water resources.
- Precipitation increased significantly from 1900 to 2005 in eastern parts of North and South America, northern Europe, and northern and central Asia. Conversely, precipitation declined in the Sahel, the Mediterranean, southern Africa, and parts of southern Asia (Figure 5).

What scientists think is likely....

- Widespread mass losses from glaciers and reductions in snow cover are projected to accelerate throughout the 21st century, reducing water availability and changing seasonality of flow patterns.
- Model projections include contraction of snow cover area, widespread increases in depth to frost in permafrost areas, and Arctic and Antarctic sea ice shrinkage.
- The incidence of extreme high sea level has likely increased at a broad range of sites worldwide since 1975.
- Based on current model simulations it is very likely that the meridional overturning circulation (MOC) of the Atlantic Ocean will slow down during the 21st century; nevertheless regional temperatures are predicted to increase. Large-scale and persistent changes in the MOC may result in changes in marine ecosystem productivity, fisheries, ocean CO₂ uptake, and terrestrial vegetation.

Figure 5. Relative changes in precipitation (in percent) for the period 2090-2099, relative to 1980-1999. Values are multi-model averages based on the SRES A1B scenario for December to February (left) and June to August (right). White areas are where less than 66% of the models agree in the sign of the change and stippled areas are where more than 90% of the models agree in the sign of the change (IPCC 2007).

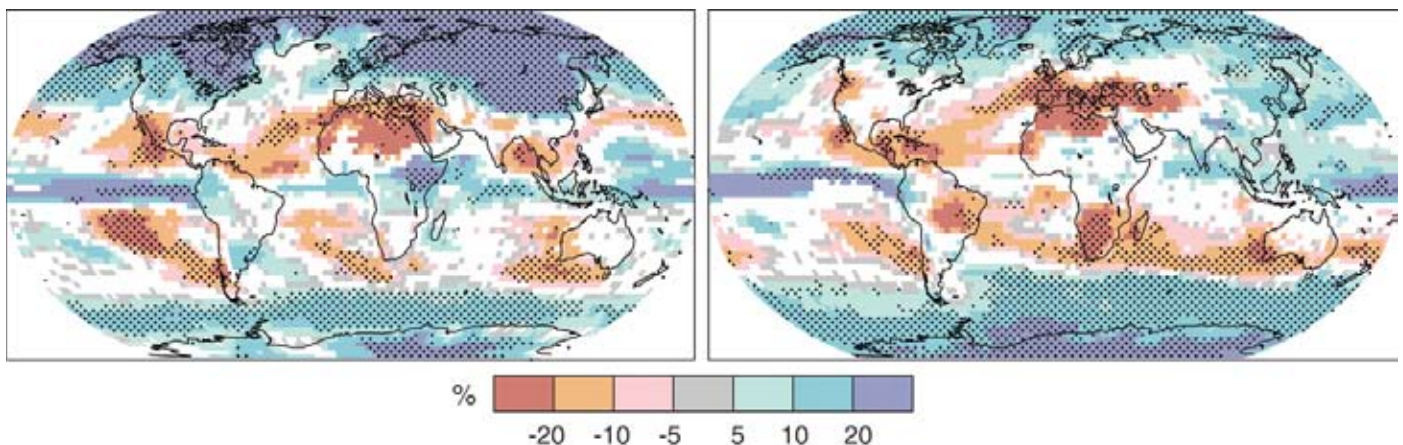


Table 2. Projected global average sea level rise at the end of the 21st century, adapted from IPCC 2007b.

Notes: a) Temperatures are assessed best estimates and likely uncertainty ranges from a hierarchy of models of varying complexity as well as observational constraints.

Emissions Scenario	Sea level rise (m at 2090-2099 relative to 1980-1999)
	Model-based range (excluding future rapid dynamical changes in ice flow)
Constant year 2000 concentrations ^a	Not available
B1 scenario	0.18 - 0.38
B2 scenario	0.20 - 0.43
A1B scenario	0.21 - 0.48
A2 scenario	0.23 - 0.51
A1F1 scenario	0.26 - 0.59

- Globally the area affected by drought has likely increased since the 1970s and the frequency of extreme precipitation events has increased over most areas.
- Future tropical cyclones (typhoons and hurricanes) are likely to become more intense, with larger peak wind speeds and increased heavy precipitation. Extra-tropical storm tracks are projected to move poleward, with consequent shifts in wind, precipitation, and temperature patterns.
- Increases in the amount of precipitation are very likely in high latitudes and decreases are likely in most subtropical land regions, continuing observed patterns (Figure 5).
- Increases in the frequency of heavy precipitation events in the coming century are very likely, resulting in potential damage to crops and property, soil erosion, surface and groundwater contamination, and increased risk of human death and injury.
- Model-based projections of global average sea level rise at the end of the 21st century range from 0.18 to 0.59 meters, depending on specific emissions scenarios (Table 2). These projections may actually underestimate future sea level rise because they do not include potential feedbacks or full effects of changes in ice sheet flow.
- Partial loss of ice sheets and/or the thermal expansion of seawater over very long time scales could result in meters of sea level rise, major changes in coastlines and inundation of low-lying areas, with greatest effects in river deltas and low-lying islands.

C. Vegetation and Wildlife

What scientists know...

- Temperature increases have affected Arctic and Antarctic ecosystems and predator species at high levels of the food web.
- Changes in water temperature, salinity, oxygen levels, circulation, and ice cover in marine and freshwater ecosystems have resulted in shifts in ranges and changes in algal, plankton, and fish abundance in high-latitude oceans; increases in algal and zooplankton abundance in high-latitude and high-altitude lakes; and range shifts and earlier fish migrations in rivers.
- A meta-analysis of climate change effects on range boundaries in Northern Hemisphere species of birds, butterflies, and alpine herbs shows an average shift

What scientists think is possible...

- Arctic late-summer sea ice may disappear almost entirely by the end of the 21st century (Figure 3).
- Current global model studies project that the Antarctic ice sheet will remain too cold for widespread surface melting and gain mass due to increased snowfall. However, net loss of ice mass could occur if dynamical ice discharge dominates the ice sheet mass balance.

- of 6.1 kilometers per decade northward (or 6.1 meters per decade upward), and a mean shift toward earlier onset of spring events (frog breeding, bird nesting, first flowering, tree budburst, and arrival of migrant butterflies and birds) of 2.3 days per decade (Parmesan and Yohe 2003).
- Poleward range shifts of individual species and expansions of warm-adapted communities have been documented on all continents and in most of the major oceans of the world (Parmesan 2006).
- Satellite observations since 1980 indicate a trend in many regions toward earlier greening of vegetation in the spring linked to longer thermal growing seasons resulting from recent warming.
- Over the past 50 years humans have changed ecosystems more rapidly and extensively than in any previous period of human history, primarily as the result of growing demands for food, fresh water, timber, fiber, and fuel. This has resulted in a substantial and largely irreversible loss of Earth's biodiversity
- Although the relationships have not been quantified, it is known that loss of intact ecosystems results in a reduction in ecosystem services (clean water, carbon sequestration, waste decomposition, crop pollination, etc.).
- Net carbon uptake by terrestrial ecosystems is likely to peak before mid-century and then weaken or reverse, amplifying climate changes. By 2100 the terrestrial biosphere is likely to become a carbon source.
- Increases in global average temperature above 1.5 to 2.5°C and concomitant atmospheric CO₂ concentrations are projected to result in major changes in ecosystem structure and function, species' ecological interactions, and species' geographical ranges. Negative consequences are projected for species biodiversity and ecosystem goods and services.
- Model projections for increased atmospheric CO₂ concentration and global temperatures significantly exceed values for at least the past 420,000 years, the period during which more extant marine organisms evolved. Under expected 21st century conditions it is likely that global warming and ocean acidification will compromise carbonate accretion, resulting in less diverse reef communities and failure of some existing carbonate reef structures. Climate changes will likely exacerbate local stresses from declining water quality and overexploitation of key species (Hoegh-Guldberg et al. 2007).
- Ecosystems likely to be significantly impacted by changing climatic conditions include:

What scientists think is likely...

- The resilience of many ecosystems is likely to be exceeded this century by an unprecedented combination of climate change, associated disturbance (flooding, drought, wildfire, insects, ocean acidification) and other global change drivers (land use change, pollution, habitat fragmentation, invasive species, resource over-exploitation) (Figure 6).
- Exceedance of ecosystem resilience may be characterized by threshold-type responses such as extinctions, disruption of ecological interactions, and major changes in ecosystem structure and disturbance regimes.
- i. terrestrial – tundra, boreal forest, and mountain regions (sensitivity to warming); Mediterranean-type ecosystems and tropical rainforests (decreased rainfall)
- ii. coastal – mangroves and salt marshes (multiple stresses)
- iii. marine – coral reefs (multiple stresses); sea-ice biomes (sensitivity to warming)
- Approximately 20% to 30% of plant and animal species assessed to date are at in-

What scientists think is possible...

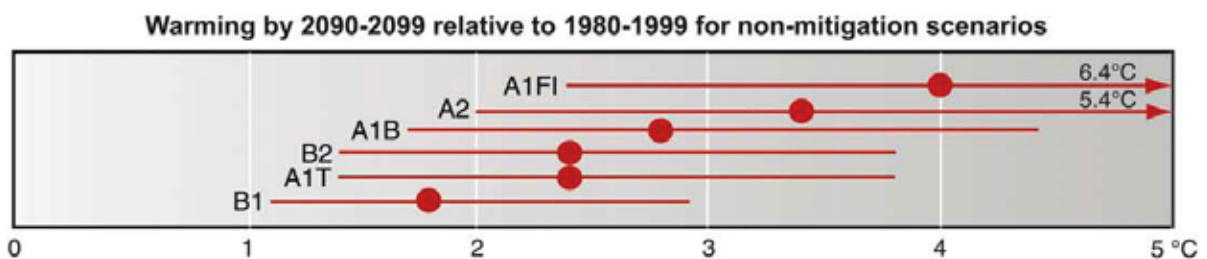
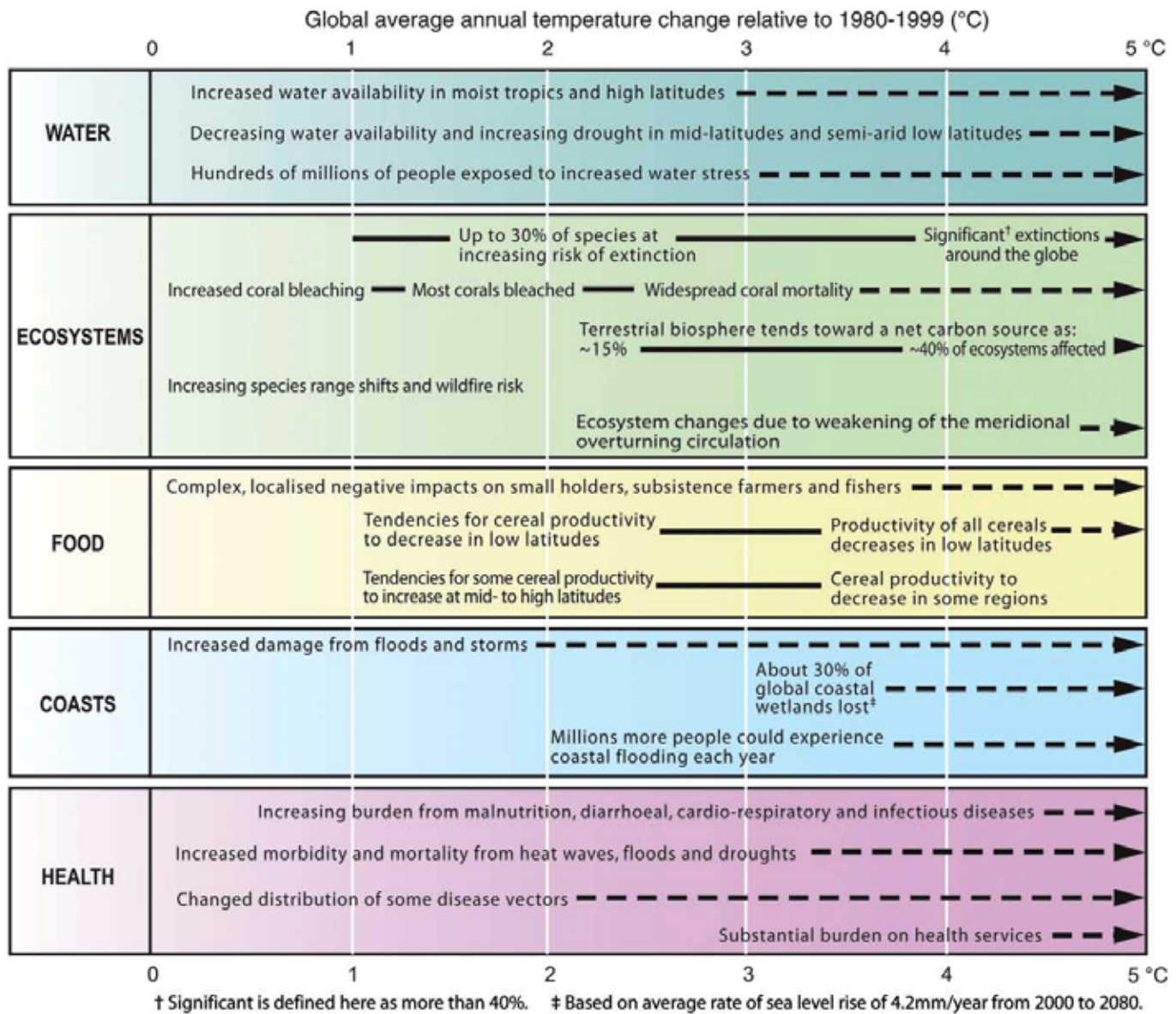


Figure 6. Examples of impacts associated with projected global average surface warming. Upper panel: Illustrative examples of global impacts projected for climate changes (and sea level and atmospheric CO₂, where relevant) associated with different amounts of increase in global average surface temperature in the 21st century. The black lines link impacts; broken-line arrows indicate impacts continuing with increasing temperature. Entries are placed so that the left-hand side of text indicates the approximate level of warming that is associated with the onset of a given impact. Quantitative entries for water scarcity and flooding represent the additional impacts of climate change relative to the conditions projected across the range of SRES scenarios A1FI, A2, B1 and B2. Adaptation to climate change is not included in these estimations. Confidence levels for all statements are high. Lower panel: Dots and bars indicate the best estimate and likely ranges of warming assessed for the six SRES marker scenarios for 2090-2099 relative to 1980-1999 (IPCC 2007a).

creased risk of extinction with increases in global average temperature in excess of 1.5 to 2.5°C.

- Endemic species may be more vulnerable to climate changes, and therefore at higher risk for extinction, because they may have evolved in locations where where paleo-climatic conditions have been stable.
- Although there is great uncertainty about how forests will respond to changing climate and increasing levels of atmospheric CO₂, the factors that are most typically predicted to influence forests are increased fire, increased drought, and greater vulnerability to insects and disease (Brown 2008b).

D. Hazards

What scientists know...

- Climate change currently contributes to the global burden of disease and premature death through exposure to extreme events and changes in water and air quality, food quality and quantity, ecosystems, agriculture, and economy (Parry et al. 2007).
- The most vulnerable industries, settlements, and societies are generally those in coastal and river flood plains, those whose economies are closely linked with climate-sensitive resources, and those in areas prone to extreme weather events.
- By the 2080s many millions more people than today are projected to experience flooding due to sea level rise, especially

those in the low-lying megadeltas of Asia and Africa and on small islands.

- Climate change affects the function and operation of existing water infrastructure and water management practices, aggravating the impacts of population growth, changing economic activity, land-use change, and urbanization.

What scientists think is likely...

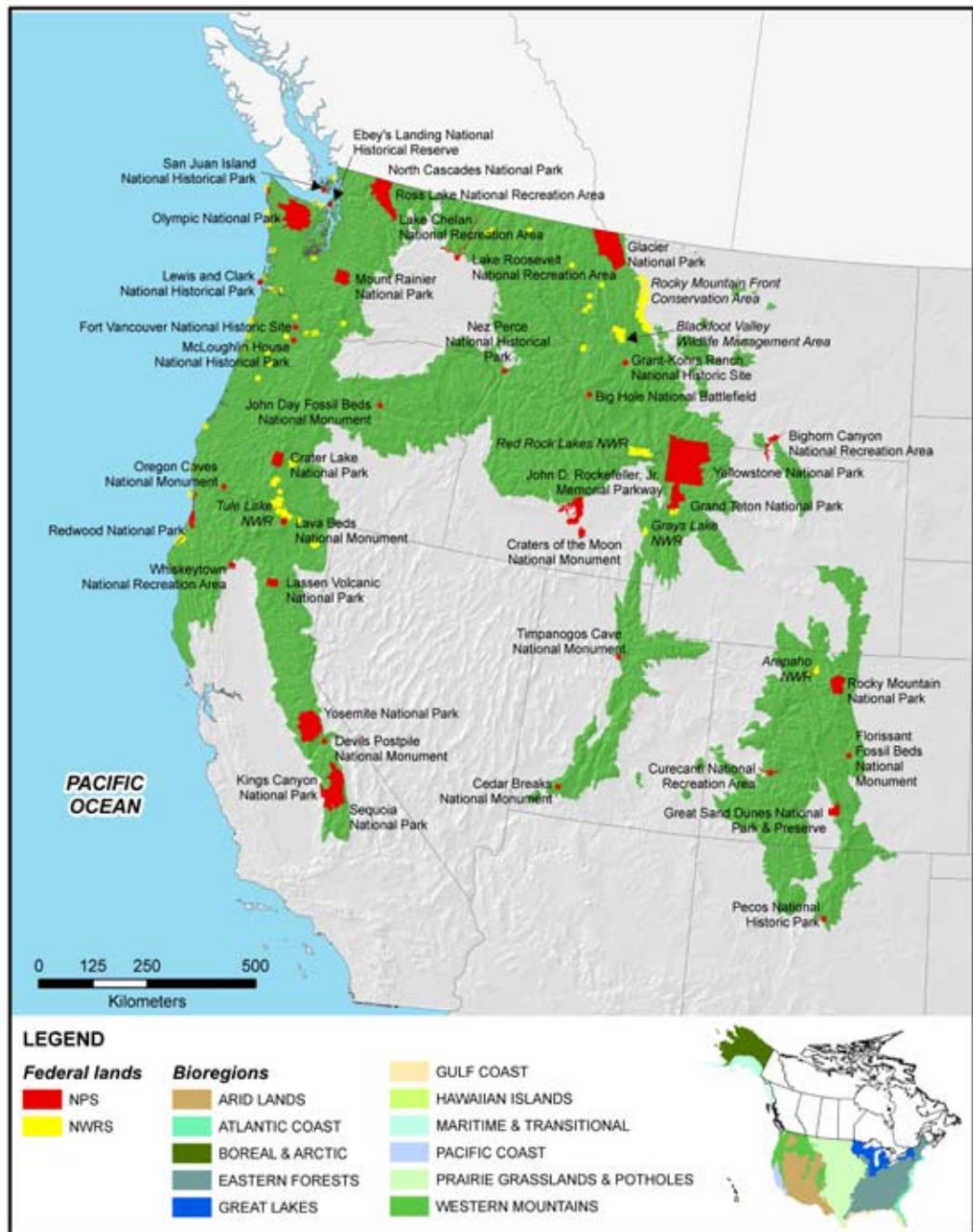
- Up to 20% of the world's population will live in areas where river flood potential could increase by the 2080s, with major consequences for human health, physical infrastructure, water quality, and resource availability.
- The health status of millions of people is projected to be affected by climate change, through increases in malnutrition; increased deaths, disease, and injury due to extreme weather events; increased burden of diarrheal diseases; increased cardio-respiratory disease due to higher concentrations of ground-level ozone in urban areas; and altered spatial distribution of vector-borne diseases.
- Risk of hunger is projected to increase at lower latitudes, especially in seasonally dry and tropical regions.

What scientists think is possible...

- Although many diseases are projected to increase in scope and incidence as the result of climate changes, lack of appropriate longitudinal data on climate change-related health impacts precludes definitive assessment.

III. Climate Change Impacts to Western Mountains and Forests

This document pertains to the Western Mountains and Forests region shown below. A list of parks and refuges for which this analysis is most useful is included in Appendix A. To help the reader navigate this section, each category is designated by color-coded tabs on the outside edge of the document.



Summary

Observed climate changes in the Western Mountains and Forests bioregion include increased seasonal, annual, minimum, and maximum temperatures, altered precipitation patterns, and a shift toward earlier timing of peak runoff. These climatic changes have resulted in widespread mortality in western forests, species range shifts and changes in phenology, productivity, and distribution of species, and an increase in wildfire severity, intensity, maximum fire size, and area burned. Predicted regional changes include an increase in average temperature of around 0.3°C per decade over the next 50 years, dramatically reduced snowpack accumulation in western mountains, and commensurate reductions in runoff and natural water storage. Ecological changes likely to result from these climatic changes include continued shifts in species phenology, productivity, and distributions, species extinctions, increased frequency, size, and duration of wildfires, increased drought length and severity, and range expansion of forest pests and pathogens.



Colorado River and surrounding river basin at Grand Canyon National Park; National Park Service photo.

A. TEMPERATURE

What scientists know....

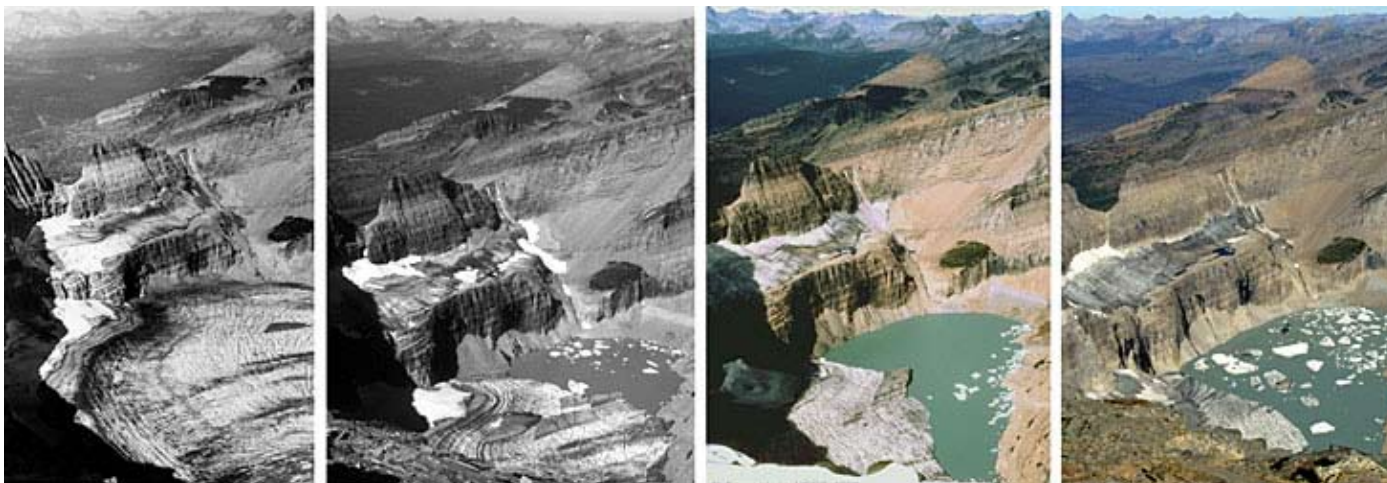
- Winter and spring temperatures increased in western North America during the twentieth century (Mote et al. 2005). The rate of change varies with location, but the central tendency is a warming of 1°C per century from 1916 to 2003 (Hamlet et al. 2007).
- The rate of temperature increase from 1947 to 2003 is roughly double that of the period from 1916 to 2003, largely attributable to the fact that much of the observed warming occurred from 1975 to the present. The largest temperature trends occurred in January-March (Hamlet and Lettenmaier 2007).
- Regionally averaged spring and summer temperatures for 1987 to 2003 were 0.87°C higher than those for 1970 to 1986, and spring and summer temperatures for 1987 to 2003 were the warmest since the beginning of the record in 1895 (Westerling et al. 2006).
- Air temperatures have increased $1.1\text{--}1.4^{\circ}\text{C}$ since the early 1990s in the Loch Vale Watershed of Rocky Mountain National Park, although this is mostly a summer phenomenon and has not influenced winter snow dynamics (Clow et al. 2003).
- In Colorado statewide temperatures have increased about 1.1°C over 30 years, with the greatest increase in the North Central Mountains and the least in southwestern Colorado, including the San Juan Mountains (Ray et al. 2008).
- Observational evidence shows that spring temperatures over western North America have undergone significant warming over the past half century, while autumn temperatures have shown relatively little change; however, recent research suggests that after accounting for seasonally opposite effects of atmospheric circulation, similar warming trends of around $+0.2^{\circ}\text{C}/\text{decade}$ exist for both seasons (Abatzoglou and Redmond 2007).

What scientists think is likely....

- Extreme cold temperatures are projected to increase faster than extreme warm temperatures during the next century (Kharin et al. 2007).
- All natural ecosystems of California are likely to be affected by changes in temperature and precipitation, including altered structure, composition, and productivity of vegetation communities, more frequent and intense wildfires, nonnative species invasions, and a significant rise in the number of threatened and endangered species (Lenihan et al. 2003).
- The average warming rate in the Pacific Northwest during the next ~50 years is expected to be in the range $0.1\text{--}0.6^{\circ}\text{C}$ per decade, with a best estimate of 0.3°C per decade. For comparison, observed warming in the second half of the 20th century was approximately 0.2°C per decade (Mote et al. 2008b).

What scientists think is possible....

- Climate model simulations for the Columbia, Sacramento/San Joaquin, and Colorado River basins over the first half of the 21st century indicate a general large-scale warming of $1\text{--}2^{\circ}\text{C}$ as compared to present (Barnett et al. 2004).
- Regional climate models indicate that on average California may experience substantially warmer and wetter winters, somewhat warmer summers, and an enhanced El Niño Southern Oscillation (ENSO) during the 21st century (Lenihan et al. 2003).
- Regional climate projections for the northwestern United States for the late 21st century include increased frequency of extreme hot events and decreased frequency of extreme cold events, and decreased severity of cold events (Diefenbaugh et al. 2005).



Repeat photography of Grinnell Glacier, Glacier National Park, showing the recession of the glacier from left to right in 1938, 1981, 1998, 2005; US Geological Survey photos.

B. THE WATER CYCLE

What scientists know....

- Between 1950 and 1999 there was a shift in the character of mountain precipitation, with more winter precipitation falling as rain instead of snow, earlier snow melt, and associated changes in river flow that include relative increases in the spring and relative decreases in the summer months (Mote et al. 2005, Barnett et al. 2008).
- The well-documented shift toward earlier peak runoff in the western United States in recent decades has been attributed to more precipitation falling as rain rather than snow and earlier snowmelt (Knowles et al. 2006).
- Historic patterns suggest that to date global warming has played a relatively minor role in determining cool season precipitation trends in the western United States. For example, trends in precipitation across the region have been generally upward since the early part of the century (owing primarily to large-scale drought in the early part of the record) but are opposite in sign for the Pacific Northwest and the Colorado River basin in the last half century (Hamlet and Lettenmaier 2007).
- Much of the mountainous west has experienced overall declines in spring snowpack since the mid-20th century, despite increases in winter precipitation in many places. The largest reductions have occurred where winter temperatures are mild, especially in the Cascade Mountains (where estimates of April 1 snow water equivalent indicate a 15–35% decline from mid-century to 2006) and in northern California. In most mountain ranges, there has been little change in snowpack at the highest elevations but major declines at lower elevation snow lines (Mote et al. 2005, Field 2007, Mote et al. 2008a).
- Snowmelt contributes 75% of all water in streams throughout the west, and acts as a water storage reservoir. Warmer wintertime temperatures and earlier melt dates will deplete this virtual reservoir, leaving much less available water for natural systems and human uses (Service 2004).
- Mountain glaciers and snow cover have declined measurably in both hemispheres (IPCC 2007a). Since 1958, North Cascades National Park, which contains over half of the glaciers in the continental United States, has experienced a 7% reduction in glacial area and a reduced mass in 80% of the park's glaciers. In 1993 the largest glaciers in Glacier National Park were measured at 72% of their 1850 areal extent and many small glaciers had vanished. The rate of glacial melting suggests that the park's remnant glaciers will be gone in the next 25 to 30 years (Hall and Fagre 2003, Burkett et al. 2005).
- Increased precipitation variability and systematic warming associated with



Flooding in North Cascades National Park in late 2006; National Park Service photo.



Yosemite National Park in the Sierra Nevada Mountains is already experiencing resource impacts from climate change; National Park Service photo.

late 20th century climate has increased flood risks in rain-dominant basins and in many near coastal areas in Washington, Oregon, and California (Hamlet and Lettenmaier 2007).

- Recent analysis suggests that the majority of the low-frequency changes in the hydrological cycle (river flow, temperature, and snow pack) observed in the western United States from 1950 to 1999 are due to human caused climate changes from greenhouse gases and aerosols, based on differences between observed and expected trends (Barnett et al. 2008).

What scientists think is likely....

- The projected large-scale warming of 1-2°C over present-day temperatures will significantly impact water resources, including a large reduction in the volume and persistence of mountain snowpacks, and a commensurate reduction in natural water storage. Current demands on water resources in many parts of the West will not be met under plausible future climate conditions, including water supplies for natural and anthropogenic systems (Barnett et al. 2004, Knowles et al. 2006).

- Warming in the western states is expected to increase the fraction of precipitation that falls as rain rather than snow and hasten the onset of snowmelt once snowpacks have formed. Snow deposition is sensitive to wintertime (November-March) warming trends, whereas snowmelt is sensitive to changes in springtime temperatures (Knowles et al. 2006).
- Reduced snowpack and earlier runoff will mean less available water for summer irrigation needs, higher water temperatures, and increased conflict between agricultural users and those whose principal concern is sustaining fish populations. These effects will be especially profound in smaller, snowmelt-driven rivers such as the Yakima River in Washington State (Barnett et al. 2004)

- Regional warming may reduce western snowpacks by up to 60% over the next

50 years in regions such as the Cascade Mountains of Oregon and Washington. Summertime streamflows are expected to decrease in response by 20 to 50% in these areas (Service 2004).

- Changes in flood risk are likely to result in substantial changes in sediment transport and channel formation processes, and are also likely to affect ecological processes that are sensitive to changes in the probability distributions of high flow events such as habitat stability, biodiversity, and trophic structure (Konrad and Booth 2005, Hamlet and Lettenmaier 2007).
- Precipitation events that are currently considered extreme (20-year return interval) are expected to occur roughly twice as often as they currently do, consistent with general increases in rain-storm intensity (Kharin et al. 2007).
- The decline of glacial ice may be linked to increases in mean summer temperature and/or a reduction in the winter snowpack that forms and maintains glaciers (Hall and Fagre 2003).

What scientists think is possible....

- Potential hydrologic and ecological responses to earlier snowmelt modeled for the Loch Vale Watershed at Rocky Mountain National Park showed 50% reductions in snowpack and 4-5 week earlier increases in soil moisture and runoff (compared to mean onset of spring conditions from 1984 to 1998) with a 4°C temperature increase and doubling of atmospheric of CO₂ (Baron et al. 2000).
- Large- and small-scale models are consistent in their predictions for future hydrologic patterns in California; these include declines in summer low streamflows (i.e. more extremely low flows in summer) and increases in winter stream flows (i.e. more winter flooding), and a shift of runoff peaks to earlier in the year (Dettinger et al. 2004, Maurer and Duffy 2005).
- The greatest reduction in snowpack volume in the Sierra Nevada region of

California is expected to occur in the elevation range of 1,300-2,700 meters. Models predict that warmer storms and earlier snowmelt will cause April snow accumulation to drop to 95% of its present levels by 2030 and 48% by 2090; these changes will result in lower spring-summer streamflows, increased salinity, and changes in species habitat in the San Francisco Estuary (the third largest estuary in the U.S.) (Knowles and Cayan 2004).

- Projected precipitation changes in the Pacific Northwest are unlikely to be distinguishable from natural variability until late in the 21st century. Most projections show winter precipitation increasing and summer precipitation decreasing and an increase in intense precipitation (Mote et al. 2008b).

C. VEGETATION

What scientists know....

- Climate has demonstrably affected terrestrial ecosystems through changes in the seasonal timing of life-cycle events (phenology), plant-growth responses (primary production), and biogeographic distribution (Parmesan 2006, Field 2007). Statistically significant shifts in Northern Hemisphere vegetation phenology, productivity, and distribution have been observed and are attributed to 20th cen-

tury climate changes (Walther et al. 2002, Parmesan and Yohe 2003, Root et al. 2003, Parmesan 2006).

- Growth of subalpine tree species in the Pacific Northwest is strongly influenced by climate factors including snowpack duration, summer moisture stress, and growing season length. Observed effects of climate changes in the region include displacement of subalpine meadows by tree species, infilling of previously open areas such as avalanche paths, and tree line changes (Fagre et al. 2003).
- The vegetation growing season, as defined by continuous frost-free air temperatures, has increased by on average about two days/decade since 1948 in the conterminous United States, with the largest changes occurring in the West (Ryan 2008).
- Increases in minimum monthly temperature and multidecadal variability in minimum temperature and precipitation are linked to increased annual branch growth of krummholz whitebark pine, invasion by whitebark pine and western white pine into formerly persistent snowfields, timing of vertical branch emergence in krummholz whitebark pine, and invasion by lodgepole pine into subalpine meadows in upper elevation forests of the central Sierra Nevada, California (Millar et al. 2004).
- Widespread mortality events during the 20th century in forests of the West have led to conditions outside ranges of variability documented in past centuries in western North America. (CIRMOUNT 2006).
- The recent drought centered on 2000–2004 caused extensive forest dieback in the West, including 1.5 million hectares of piñon pine and 1.0 million hectares of ponderosa pine on the Colorado Plateau. Bark beetles were frequently the ultimate cause of death in dense and climate-stressed forests of this region (CIRMOUNT 2006).

Hidden Meadow Camp at North Cascades National Park is losing ground as trees encroach into areas where they were previously unable to survive. National Park Service photo.





Whitebark pine along the caldera rim at Crater Lake National Park; National Park Service photo.

- Forests in the western United States sequester 20 to 40% of U.S. carbon emissions from fossil fuels, an amount equal to about one-half of the carbon absorbed by terrestrial ecosystems in the conterminous U.S. (Pacala et al. 2001, Westering et al. 2006, Brown 2008). Changes in forest productivity, structure, and composition will result in changes in the rate of carbon sequestration and amount of carbon stored as biomass.

- Some aquatic and terrestrial plant and microbial communities are significantly altered by nitrogen deposition, either from low level atmospheric deposition or “hotspot” deposition downwind of metropolitan areas or agricultural operations. Observed vegetation responses include increased plant productivity, biotic community changes, and deleterious effects on sensitive organisms including lichens and phytoplankton (Fenn et al. 2003).

What scientists think is likely....

- Vegetation response to changing climate will depend on the factors that limit productivity at a particular site; for example, changes in growing season length may affect annual productivity, and increased nitrogen and CO₂ inputs strongly influence forest productivity if other factors (water, temperature, radiation) are less limiting (Ryan 2008).
- Forest growth is expected to decrease in regions where increased temperature coincides with decreased precipitation (western Alaska, Interior West, Southwest) (Ryan 2008).
- Rising CO₂ will very likely increase photosynthesis for forests, but the increased photosynthesis will likely only increase wood production in young forests on fertile soils. Where nutrients are not limiting, rising CO₂ increases photosynthesis and wood production, but on infertile soils the extra carbon from increased photosynthesis will be quickly respired. The response of older forests to CO₂ is uncertain, but possibly will be lower

than the average of the studied younger forests (Ryan 2008).

- For California, models indicate a shift in dominance from needle-leaved toward broad-leaved trees, along with increases in vegetation productivity, especially in the relatively cool and mesic (moderately moist habitat, such as temperate hardwood) regions of the state (Lenihan et al. 2003).

What scientists think is possible....

- Future climate change scenarios for California predict a decline in alpine/subalpine forest cover, increases in the productivity of evergreen hardwoods and the subsequent displacement of evergreen conifer forest by mixed evergreen forest, and expansion of grasslands (Lenihan et al. 2008).
- Warmer winter temperatures may reduce the competitiveness of Douglas-fir seedlings by not providing a sufficient period of winter cold temperatures to meet the species' chilling requirements; furthermore, the distribution of Douglas-fir may decline if fire frequencies exceed 1 fire per 20 years (Whitlock et al. 2003).
- Dramatic changes in forested lands within the western U.S are expected to occur with climate shifts; these include severe contraction and northward displacement of alpine habitats, subalpine spruce-fir forests, and aspen; fragmentation and disjunction of species' ranges; and possible expansion of forests into current grass- and shrublands (ISAB 2007).
- Douglas fir, Pacific yew and red alder are predicted to shift in range from west of to east of the Cascades, and the potential habitats of dominant rainforest conifers (e.g., western hemlock, western red cedar) are expected to decrease west of the Cascades but expand into mountain ranges of the interior West. In contrast, Ponderosa pine, which is tolerant of rela-



View up into old growth Douglas-fir canopy at Olympic National Park; National Park Service photo.



Edith's checkerspot butterfly at Pinnacles National Monument; National Park Service photo.

tive warm and dry climate, is predicted to expand its range significantly and to occur west of the Cascades (ISAB 2007).

D. WILDLIFE

What scientists know....

- A consistent temperature-related shift has been observed across a broad range of plant and animal species (80% of species from 143 studies), including changes in species density, north- or poleward range shifts, changes in phenology, and shifts in genetic frequencies (Root et al. 2003).
- A meta-analysis of climate change effects on range boundaries in Northern Hemisphere species of birds, butterflies, and alpine herbs shows an average shift of 6.1 kilometers per decade northward (or meters per decade upward), and a mean shift toward earlier onset of spring events (frog breeding, bird nesting, first flowering, tree budburst, and arrival of migrant butterflies and birds) of 2.3 days per decade (Parmesan and Yohe 2003).
- Many North American wildlife species have shifted their ranges to more northerly or higher elevations. For example, the Edith's checkerspot butterfly has become locally extinct in the southern, low-elevation portion of its North American range, but has extended its range 90 kilometers north and 120 kilometers upward in elevation (Field 2007).
- The northern boundary of the sache skipper butterfly has expanded 670 kilometers from California to Washington State in the past 40 years, corresponding with areas where the winter minimum temperature has rise 3°C since 1950 (Crozier 2003). During a single year -1998, the second warmest on record - the butterfly moved 120 kilometers northward, and laboratory and field manipulations show that because winter cold extremes dictate the northern range limit, the butterfly may extend its range even further as minimum temperatures increase (Crozier 2003, 2004).

- Past physiological studies suggest that stress from increasing temperatures is the likely cause of rapid population decline and local extinctions of American pika, a cold-adapted, mountain-dwelling species that survives in isolated mountaintop islands throughout the western United States (Beever et al. 2003, NRC 2008)
- A small-mammal survey in Yosemite National Park showed substantial upward changes (~500 meters) in elevation limits for half of 28 species monitored. In addition, high-elevation species experienced range contractions and low-elevation species expanded their ranges upward (Moritz et al. 2008).

What scientists think is likely....

- American robins in the Colorado Rocky Mountains are arriving earlier in spring than in previous decades (14 days earlier in 1999 than in 1981). The interval between arrival date and the first date of bare ground has grown by 18 days, because the growing season onset date is occurring earlier in the year at low elevations but has remained largely unchanged at high elevation sites (Inouye et al. 2000).
- Changes in minimum and mean temperatures and amount and timing of precipitation may increase prevalence of hantavirus, plague, and West Nile virus in wildlife populations though changes in phenological development, increased rates of reproduction and survival, altered geographic distributions, and expansion of favorable habitats of disease vector, host, and reservoir species (Patz et al. 2000, Epstein 2001, Field 2007).

What scientists think is possible....

- Modeling of potential future climate and vegetation scenarios indicates that potential Canada Lynx habitat may decrease significantly by 2100 because of reduced snow cover and vegetation changes. Areas that may become unsuitable for lynx include the Bridger-Teton (Wyoming), Idaho Panhandle (Idaho),



American Robin; Fish and Wildlife Service photo.



Stages of life of Chinook salmon. From top to bottom: Egg, alevin, fry, smolt, ocean adult, and spawning stage; Fish and Wildlife Service images.

Kootenai (Montana), Okanogan (Washington) and Wenatchee (Washington) National Forests; high-altitude areas in Colorado; and Yellowstone and Grand Teton National Parks (Wyoming) (Gonzalez et al. 2007).

- An analysis of potential climate change impacts on mammalian species in U.S. national parks indicates that on average about 8% of current mammalian species diversity may be lost. The greatest losses across all parks occurred in rodent species (44%), bats (22%), and carnivores (19%) (Burns et al. 2003).

- Fragmentation of large ecosystems may occur due to increased disturbance and vegetative change, disrupting existing wildlife ranges. Certain types of habitat, such as margin or edge areas, will be particularly sensitive (McCarty 2001).

- Climate change has the potential to affect most freshwater life history stages of trout and salmon. Increased frequency and severity of flood flows during winter can affect over-wintering juvenile fish and incubating eggs in the streambed. Eggs of fall and winter spawning fish, including Chinook, coho, chum, and sockeye salmon and bull trout, may suffer higher levels of mortality when exposed to increased flood flows. Higher winter water temperatures could also accelerate embryo development and cause premature emergence of fry (ISAB 2007).

- Models predict that changes in stream temperatures resulting from warming of average air temperature will result in earlier spawning for greenback cutthroat trout in the central Rocky Mountains. (Cooney et al. 2005).

- Changing vegetation cover in many park areas will affect wildlife species dependent on those habitats. Animals will eventually occupy landscapes vacated by glacial ice, and utilize new alpine lakes after ice is gone (Burkett et al. 2005).

- The synergism of rapid temperature rise and stresses such as habitat destruction may disrupt connectedness among spe-

cies, lead to reformulation of species communities, and result in numerous extirpations and/or extinctions (Root et al. 2003).

- Modeled effects of elevational limits on the extinction risk of landbirds (87% of all bird species) projected a best guess of 400–550 landbird extinctions, and approximately 2150 additional species at risk of extinction, by 2100. For Western Hemisphere landbirds, intermediate extinction estimates based on climate-induced changes in actual distributions ranged from 1.3% (for 1.1°C warming) to 30.0% (for 6.4°C warming) of these species (Sekercioglu et al. 2007).

E. DISTURBANCE

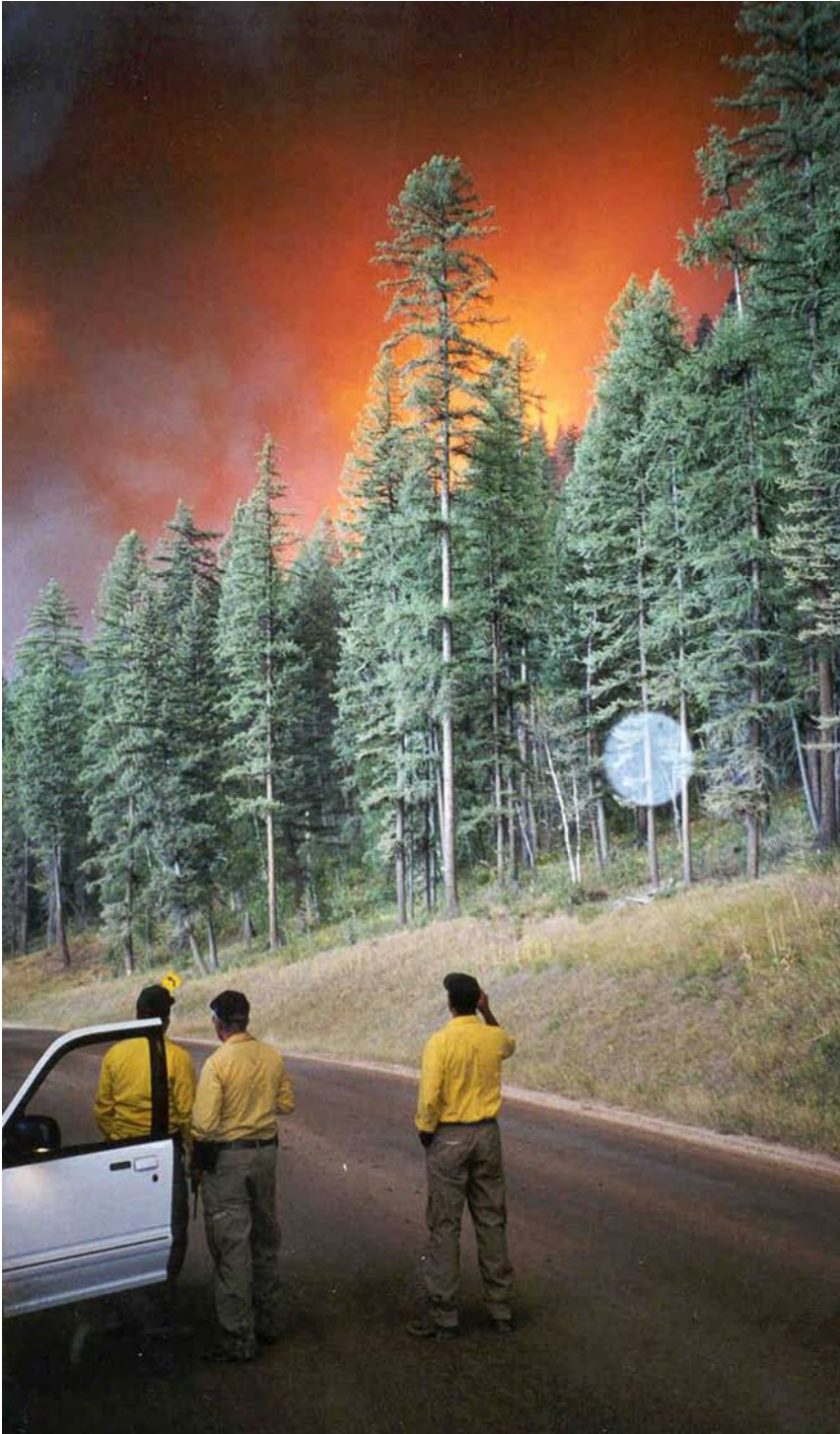
What scientists know....

- The forested area burned in the western U.S. from 1987 to 2003 was more than six and a half times the area burned from 1970 to 1986. Other observed trends include more frequent large wildfires (greater than 400 ha in size), longer wildfire durations, and longer wildfire seasons. The greatest increases in wildfire activity occurred in mid-elevation Northern Rockies forests (Westerling et al. 2006).
- A large area (approximately 120,000 km²) of California and western Nevada experienced a notable increase in the extent of forest stand-replacing (high severity) fire and increases in mean and maximum fire size and area burned annually between 1984 and 2006. These changes in fire severity and activity are attributed to a regional increase in temperature and a long-term increase in annual precipitation (Miller et al. 2008).
- Climate is a strong driver of 20th century fire synchrony in the interior west of North America. In both the inland Northwest (interior Oregon, Washington, and southern British Columbia) and in the northern Rocky Mountains warm spring–summers and warm-dry summers are associated with widespread fires. Spring climate likely affected the length of the fire season via the effects of snow-

melt on soil and fuel moisture, whereas summer climate influenced fuel moisture during the fire season (Heyerdahl et al. 2008, Morgan et al. 2008).

- More than 55 million acres of forested lands are currently impacted by disturbance. The largest agents of change are insects and pathogens, which cause an estimated financial loss of 3.7 billion dollars per year (Dale et al. 2001).

The Robert Fire was one of three large fires that burned a total of 18% of Glacier National Park's vegetation cover in 2003; National Park Service photo.



- A recent study in British Columbia linked an increase (at an increasing rate) in the number of mountain pine beetle infestations since 1970 with an increase in the amount of available optimal habitat. The amount of climatically optimal habitat has increased by more than 75% since around 1970, probably as the result of an average annual temperature increase of $>1^{\circ}\text{C}$ (Carroll 2006).
- Cumulative stresses in western mountain forests, triggered by climate change and fire exclusion, are affecting the resilience of forests. When periodic multiyear droughts, common in western North America, occur against a backdrop of rising temperatures and unnaturally high fuel loads, forests are vulnerable to insect and disease epidemics and severe fires. These combined effects can cause significant changes in forest composition and structure and extensive forest diebacks (CIRMOUNT 2006).

What scientists think is likely....

- Disturbance may reset and rejuvenate some ecosystems in some cases, and cause enduring change in others. For example, drought may weaken trees and make them susceptible to insect attack and death; however, fire is an integral component of many forest ecosystems and many tree species (e.g. lodgepole pine forests) depend on fire for regeneration (Ryan 2008).
- Disturbance events in the future may be larger and more common than those experienced historically, and planning for disturbances should be encouraged (Dale et al. 2001).
- Increases in frequency, size, and duration of wildfires in the western U.S. observed within the past three decades are attributed to a 78-day increase in the length of the wildfire season, increases in spring-summer temperatures of 0.87°C , and earlier spring snowmelt (Westerling et al. 2006).
- An increase in the average length and intensity of summer drought in the



In Great Basin National Park, fir engraver beetles leave trails beneath the bark of living and dead conifers; National Park Service photo.

Northern Rockies and elsewhere in the western United States will result in an increased frequency of large wildfires and subsequent changes in forest composition. Reduced tree densities associated with large and severe wildfires may cause western forests to become a source of increased atmospheric carbon dioxide rather than a sink, even under a relatively modest temperature-increase scenario (Westerling et al. 2006).

- Over the past century, and at the scale of the whole western United States and Canada, variations and trends in temperature have been significantly correlated with numbers of large fires and areas burned. This suggests that regional to continental warming patterns are beginning to influence fire activity at the broadest scales (CIRMOUNT 2006).
- By 2070 the length of the fire season could be increased by two to three weeks in the northern Rockies, Great Basin, and Southwest as the result of increases in summer temperature and decreases in summer humidity (Barnett et al. 2004).
- Large outbreaks of forest insects are likely influenced by observed increases in temperature because temperature controls life cycle development rates, influences synchronization of mass attacks required to overcome tree defenses, and determines winter mortality rates. Climate also affects insect populations

indirectly through effects on hosts; for example, drought stress reduces the ability of a tree to mount a defense against insect attacks (Logan and Powell 2001, Ryan 2008).

- Future northward range expansion attributed to warming temperatures has been predicted for mountain pine beetle, including potential invasion of jack pine, a suitable host that extends across the boreal forest of North America (Logan and Powell 2001).
- Resilience to perturbations, establishment of invasive species, and local extinction depend in large part on the degree of climatic stress that occurs directly before and after large ecological disturbances (CIRMOUNT 2006).

What scientists think is possible....

- Predicted trends in climate will reinforce the tendency toward longer fire seasons and accentuate conditions favorable to the occurrence of large, intense wildfires (e.g. increased storminess, higher fuel loads, drier conditions late in the season) (Lenihan et al. 2003, Whitlock et al. 2003, Westerling et al. 2006).
- Future climate change scenarios for California predict an increase in total area burned of 9-15% above the historical norm by 2100 (Lenihan et al. 2008).
- Simulations of potential future climate and vegetation indicate that future fire conditions in some parts of the northwestern US could be more severe than they are today. Even if summer precipitation in the region increased significantly, greater evapotranspiration as a result of higher temperatures could still lead to increased drought stress. Changes in temperature and the timing of precipitation would also affect fuel moisture levels, with drier fuels increasing the potential for fires (Whitlock et al. 2003).
- Modeling indicates that reduced net primary productivity and altered disturbance patterns may be expected in dry, east-side forest ecosystems in Montana

Lodgepole pine seedlings regenerate after fire in Yellowstone National Park; National Park Service photo.



and Washington under climatic warming conditions (Fagre and Peterson 2000).

F. VISITOR EXPERIENCE

- A visitor survey at Rocky Mountain National Park suggest that while a small number of visitors may increase park visits under a warming climate, more than 70% of respondents indicated that current opportunities for viewing conifer forests and wildflowers were important reasons for visitation to the park (Richardson and Loomis 2004).
- Surveys on hiking trails within National Forests in Colorado and Montana indicate that crown fires reduce the value of recreation (measured as changes in the annual value of trips taken) because of reduced aesthetic value of the landscape as perceived by respondents (Hesseln et al. 2004).
- In surveys of visitors to National Forests in Colorado, years since a non-crown fire increased the trip demand of hikers while the presence of crown fires reduced the number of mountain bike trips within the forests (Loomis et al. 2001).
- The locations of climatically ideal tourism conditions are likely to shift toward higher latitudes under projected climate change, and as a consequence spatial and temporal redistribution of tourism

A hiker enjoys Ranger Meadow in Deadman Canyon, in the southern part of Kings Canyon National Park; National Park Service photo.



activities may occur. The effects of these changes will depend greatly on the flexibility demonstrated by institutions and tourists as they react to climate change (Amelung et al. 2007).

- A statistical model of monthly visitation and anticipated climate and environmental change at Waterton Lakes National Park in the Canadian Rockies projected that annual visitation would increase between 6% and 10% in the 2020s and between 10% and 36% in the 2050s as the result of an increase in the length and quality of warm-weather tourism seasons. However, 56% of respondents indicated that they would reduce or discontinue visits to the park under projected 2080s environmental conditions, which included a decline in wildlife species, disappearance of glaciers from the park, loss of rare plant species, transition from forests to grasslands, and large increase in forest fires, (Scott et al. 2007).
- Changes to the terrestrial and aquatic species compositions in parks and refuges are likely to occur as ranges shift, contract, or expand (Burns et al. 2003).
- Parks and refuges may not be able to meet their mandate of protecting current species within their boundaries, or in the case of some refuges, the species for whose habitat protection they were designed. While wildlife may be able to move northward or to higher elevations to escape some effects of climate change, federal boundaries are static (Burns et al. 2003).
- Lakeshore levels and stream depths will increasingly fluctuate, potentially making fixed docks and boat ramps unusable for much of the year. Navigational hazards and new sand bars may be exposed (Scavia et al. 2002).
- Effects of projected climate changes on human health include increased incidence of heat stress and heat stroke, respiratory distress from pollutants released during wildfires, cardio-respiratory morbidity and mortality associated with ground-level ozone, and injury and

Temperature

The Water Cycle

Vegetation

Wildlife

Disturbance



Sprague Lake at Rocky Mountain National Park; National Park Service photo.

death from floods, storms, fires, and droughts (Epstein 2001, Confalonieri et al. 2007).

- Glaciers in western North America are important tourist attractions for mountain parks. Observed and projected changes in glacial extent (Hall and Fagre 2003) may have a negative affect on the number of visitors to Glacier National Park (Scott et al. 2007).
- Because of increases in fire season length and severity it is possible that visitors to mountain parks may experience more restrictions on their activities (e.g. campfire bans; trail and park closures) (Scott et al. 2007).
- Climate changes may favor zoonotic disease transmission to humans through altered distributions of pathogens and disease vectors, increased populations of reservoir or host species, and increased prevalence of diseases within host and reservoir populations. Diseases likely to increase in scope and/or incidence in the region include hantavirus pulmonary syndrome, plague, and West Nile virus (Epstein 2001, Confalonieri et al. 2007).
- Increasing frequency and intensity of severe storms and floods may pose threats to historic structures, roads and trails, archeological sites, administrative facilities, and other park resources and infrastructure.
- Increased summer temperatures will lead to increased utility expenditures in parks in the summer and, potentially, decreases in the winter.
- Potentially poorer visibility due to smoke from increased wildland fire activity will likely cause a negative impact on visitor experiences.

IV. No Regrets Actions: How Parks, Refuges, and Their Partners Can Do Their Part

Educate staff and the public

- Incorporate climate change research and information in interpretive and education outreach programming.
- Post climate change information on informational bulletin boards and websites.
- Educate park and refuge employees, partners, and visitors about climate change and impacts to resources through training, displays, and by offering information for further exploration and knowledge.
- Distribute up-to-date interpretive products (e.g., the National Park Service-wide Climate Change in National Parks brochure) and support the development of region, park, or refuge-specific interpretive products on the impacts of climate change.
- Develop climate change presentations for local civic organizations, user and partner conferences, etc.
- Incorporate climate change questions and answers in park-based Junior Ranger programs.
- Help visitors make the connection between reducing greenhouse gas emissions and resource stewardship.
- Encourage visitors to use public or non-motorized transportation to and around parks.

An interpretive brochure about climate change impacts to National Parks was created in 2006 and was distributed widely.

Climate Change in National Parks



- Encourage visitors to reduce their carbon footprint in their daily lives and as part of their tourism experience (For more information see http://www.epa.gov/climatechange/emissions/ind_calculator.html)

Improve sustainability and energy efficiency

- Use more energy efficient products, such as energy efficient light bulbs and ENERGY STAR approved office equipment.
- Initiate an energy efficiency program to monitor energy use in buildings, and offer guidelines for reducing energy consumption.
- Convert to renewable energy sources such as solar or wind generated power.
- Specify “green” designs for construction of new or re-modeled buildings.
- Provide a shuttle service for visitor travel to and within the park.
- Include discussions of climate change in the park Environmental Management System.
- Request and hold Climate Friendly Park workshops in cooperation with the EPA.
- Provide alternative transportation options for within-park commuting, such as employee bicycles and shuttles.
- Provide incentives for use of alternative transportation methods for park visitors and employees.
- Provide hybrid electric or propane-fueled vehicles for official use, and impose fuel standards for park vehicles. Reduce the number and/or size of park vehicles or boats to maximize efficiency.
- Provide recycling options for solid waste and trash that is generated at the park.



At the National Mall, Park Service employees use clean-energy transportation to lead tours; National Park Service photo.



Solar panels and propane generator at Yellowstone National Park; National Park Service photo.

- Use teleconferences in place of travel to conferences and meetings.
- Establish an in-park sustainability team and develop sustainability Best Management Practices.

Plan ahead and collaborate with others

- Develop strategies for managing uncertainty surrounding climate change effects in parks.
- Build a strong partnership-based foundation for future conservation efforts.
- Identify strategic priorities for climate change efforts when working with partners.
- Incorporate anticipated climate change impacts, such as decreases in lake levels and changes in vegetation and wildlife, into management plans.
- Incorporate products and services that address climate change in the development of Comprehensive Interpretive Plans
- Encourage research and scientific study in the park units and refuges.
- Engage and enlist collaborator support (e.g., tribes, nearby agencies, private landholders) in climate change discussions, responses, and mitigation.
- Take inventory of the facilities/boundaries/species within your park or refuge that require climate change mitigation or adaptation.

- Participate in gateway community sustainability efforts.

Restore damaged landscapes

- Restoration efforts are important as a means for enhancing species' ability to cope with stresses and adapt to climatic and environmental changes.
- Through restoration of natural areas, we can lessen climate change impacts on species and their habitats. These efforts will help preserve biodiversity, natural resources, and recreational opportunities.
- Restore and preserve connectivity within habitats, protect and enhance instream flows for fish, and maintain and develop access corridors to climate change refugia.
- Strategically focus restoration efforts both in terms of the types of restoration undertaken and their national, regional, and local scale and focus to help maximize resources.
- Restoration of natural hydrologic functions of coastal wetlands can be used to help protect coastal areas against potential hurricanes and flooding.

The Climate Friendly Parks Program is a joint partnership between the U.S. Environmental Protection Agency and the National Park Service. Climate Friendly Parks from around the country are leading the way in the effort to protect our parks' natural and cultural resources and ensure their preservation for future generations; National Park Service image.



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VI. Appendix A

This analysis is most useful for the following U.S. national park units and national wildlife refuges in the Western Mountains and Forests region.

U.S. National Park Units

- Aztec Ruins NM*
- Bandelier NM
- Black Canyon of the Gunnison NP
- Bryce Canyon NP
- Capulin Volcano NM
- Cedar Breaks NM
- City of Rocks NR
- Crater Lake NP
- Devils Postpile NM
- Devils Tower NM
- Florissant Fossil Beds NM
- Glacier NP
- Grand Canyon NP
- Grand Teton NP
- Great Basin NP
- Guadalupe Mountains NP
- Jewel Cave NM
- Kings Canyon NP
- Lake Chelan NRA
- Lake Roosevelt NRA
- Lassen Volcanic NP
- Lava Beds NM
- Mesa Verde NP
- Mount Rainier NP
- Mount Rushmore NM
- Natural Bridges NM
- Nez Perce NHP
- North Cascades NP
- Pinnacles NM
- Rocky Mountain NP
- Ross Lake NRA
- Olympic NP
- Oregon Caves NM
- Santa Monica Mountains NRA
- Sequoia NP
- Timpanogos Cave NM
- Walnut Canyon NM
- Whiskeytown-Shasta-Trinity NRA
- Wind Cave NP
- Yellowstone NP
- Yosemite NP
- Zion NP

National Wildlife Refuges

- Ankeny NWR
- Baskett Slough NWR
- Bear Valley NWR
- Benton Lake NWR
- Blackfoot Valley WMA
- Blue Ridge NWR
- Blue Ridge NWR
- Butte Sink WMA
- Camas NWR
- Clear Lake NWR
- Cold Springs NWR
- Columbia NWR
- Colusa NWR
- Conboy Lake NWR
- Deer Flat NWR
- Delevan NWR
- Dungeness NWR
- Franz Lake NWR
- Grasslands WMA
- Hart Mountain NWR
- Julia Butler Hansen NWR
- Kern NWR
- Klamath Marsh NWR
- Kootenai NWR
- Lee Metcalf NWR
- Little Pend Oreille NWR
- Lost Trail NWR
- Lower Klamath NWR
- Malheur NWR
- McKay Creek NWR
- McNary NWR

- Merced NWR
- Modoc NWR
- National Bison Range
- Nine-Pipe NWR
- Nisqually NWR
- North Central Valley WMA
- Pablo NWR
- Pierce NWR
- Pixley NWR
- Protection Island NWR
- Red Rock Lakes NWR
- Ridgefield NWR
- Rocky Mountain Front Conservation Area
- Sacramento NWR
- Sacramento River NWR
- San Juan Islands NWR
- San Luis NWR
- Steigerwald Lake NWR
- Stone Lakes NWR
- Sutter NWR
- Swan River NWR
- Toppenish NWR
- Tualatin River NWR
- Tule Lake NWR
- Turnbull NWR
- Umatilla NWR
- Upper Klamath NWR
- William L. Finley NWR
- Willow Creek-Lurline WMA

*Acronym	Unit Type
NHP	National Historic Park
NM	National Monument
NP	National Park
NRA	National Recreation Area
NWR	National Wildlife Refuge
WMA	Wildlife Management Area

The Department of the Interior protects and manages the nation's natural resources and cultural heritage; provides scientific and other information about those resources; and honors its special responsibilities to American Indians, Alaska Natives, and affiliated Island Communities.

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